

## Lecture summary for 21 February 2018 by Andrew Kieckhefer

Related Reading: Cushman

### Definitions/Equations:

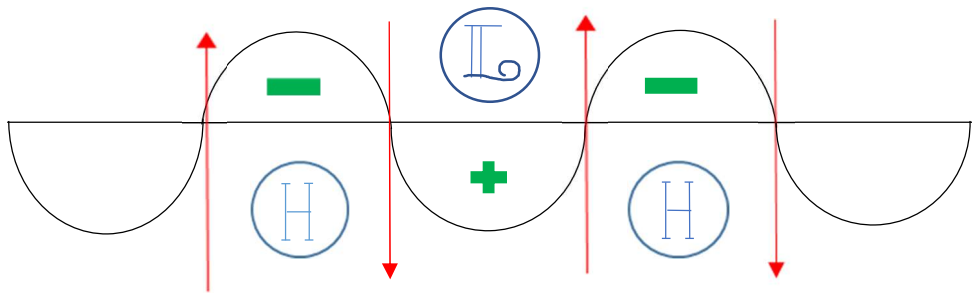
Shallow Water System: Fluids, atmospheric or oceanic, composed of a horizontal length scale that is longer than their depth scale.

$\Psi = \frac{-A}{k^2} \sin kx$  where  $A$  = amplitude and  $k$  = scale effect.  $\frac{-A}{k^2}$  produces a cosine looking wave.

$k = \frac{2\pi}{L_x}$  where  $L_x$  = wavelength

### Summary:

When graphing our stream function, we must find relations in minima/maxima and relate our stream function to wind and vorticity. Given a stream function where  $A > 0$  we look at the following wave.



The red arrows depict meridional winds, alternating in orientation. L indicates a low in the stream function while H indicates a high achieved in the stream function. Vorticity maxima and minima are depicted with + and - respectively. They depend on the size of the amplitude ( $A$ ).

The space scale ( $k$ ) relates to our wavelength ( $L_x$ ). When the space scale ( $k$ ) is large, the stream function wavelength ( $L_x$ ) and amplitude are small. In this case the stream function amplitude ( $A$ ) is small. This relation results in a small disturbance with small stream function abnormalities. The opposite occurs with a small space scale ( $k$ ). A large wavelength ( $L_x$ ) results in stronger meridional wind velocities.

In a shallow water system hydrostatic balance holds, shown by scaling  $\frac{H}{L} < 1$ . The divergence, represented by  $u = u(x, y)$  and  $v = v(x, y)$  does not reference depth, **maintaining 2D divergence**. Proven further,  $\frac{\partial u}{\partial z} = \frac{\partial v}{\partial z} = 0$ . The system is boussinesq, so  $\rho = \rho_0$ . Friction may be ignored,  $Ek \ll 1$  while  $R_0$  and  $R_{0T} \leq 1$ . The Rossby numbers may be **less than or equal to one**, flow is not perfectly geostrophic.